

UW-Madison Efforts as Part of the HFIP HRH Test

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Outline

- Overview of Model Physics, Dynamics and Configuration
- Results and Lessons Learned
(aka Trials, Tribulations and ... Simulations)
- Future Plans

UW-NMS Dynamics

$$\frac{\partial u_i}{\partial t} = -\varepsilon_{i,j,k} (\zeta_j + f_j) u_k - \frac{\partial k}{\partial x_i} - \theta \frac{\partial \pi}{\partial x_i} + \delta_{i3} g$$

$$\varepsilon_{i,j,k} = \begin{cases} 0 & \text{if } i=j, \text{ or } j=k, \text{ or } i=k \\ 1 & \text{if } i,j,k \text{ are } 1,2,3 \text{ or } 2,3,1 \text{ or } 3,1,2 \\ -1 & \text{if } i,j,k \text{ are } 3,2,1 \text{ or } 2,1,3 \text{ or } 1,3,2 \end{cases}$$

$$\pi = c_p \left(\frac{p}{1000mb} \right)^{\frac{R}{c_p}}$$

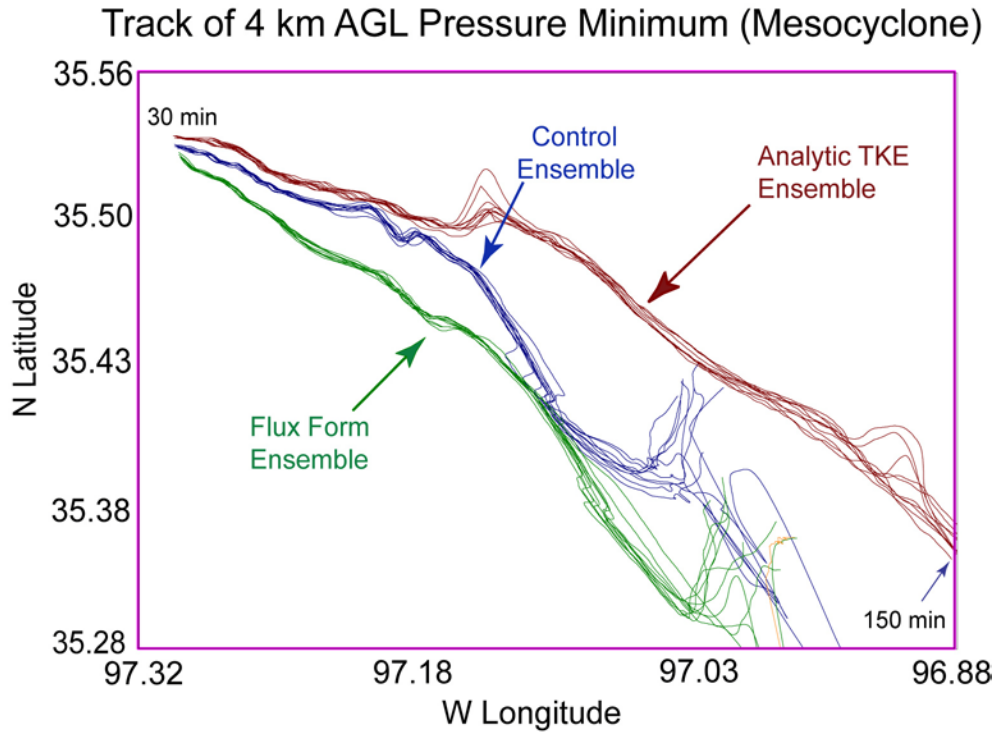
$$k = \frac{1}{2} (u_i^2)$$

$$\zeta_i = \varepsilon_{i,j,k} \frac{\partial u_j}{\partial x_k}$$

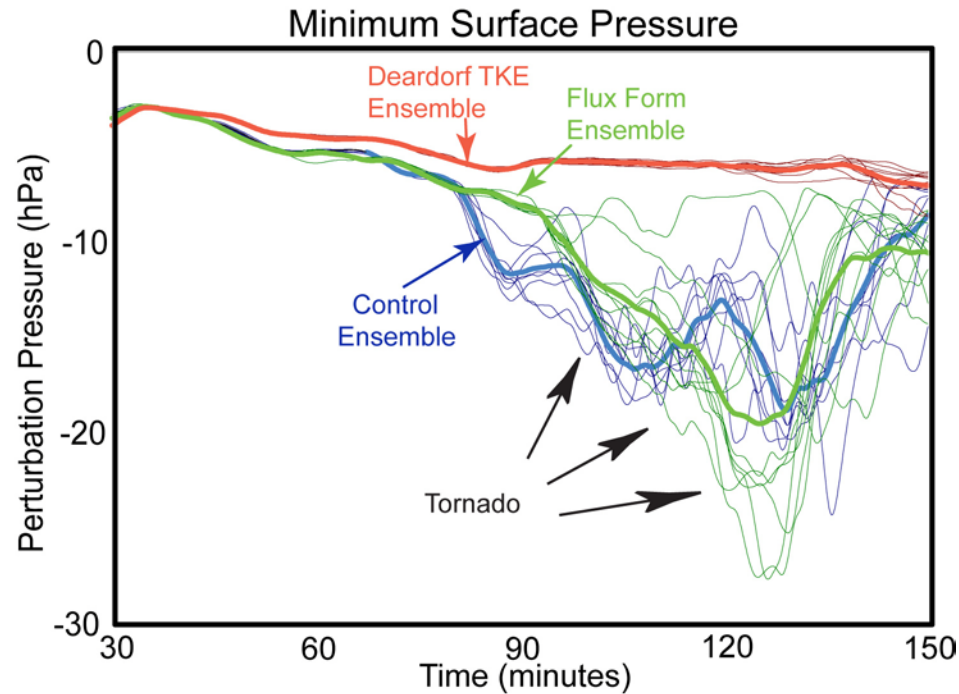
$$\varepsilon_{i,j,k} A_j B_k = \vec{A} \times \vec{B}$$

Finite differencing accomplished with Arakawa and Lamb (1981) vorticity, kinetic energy and enstrophy conserving scheme.

Conservative Dynamics Impact



Conservative Dynamics Impact



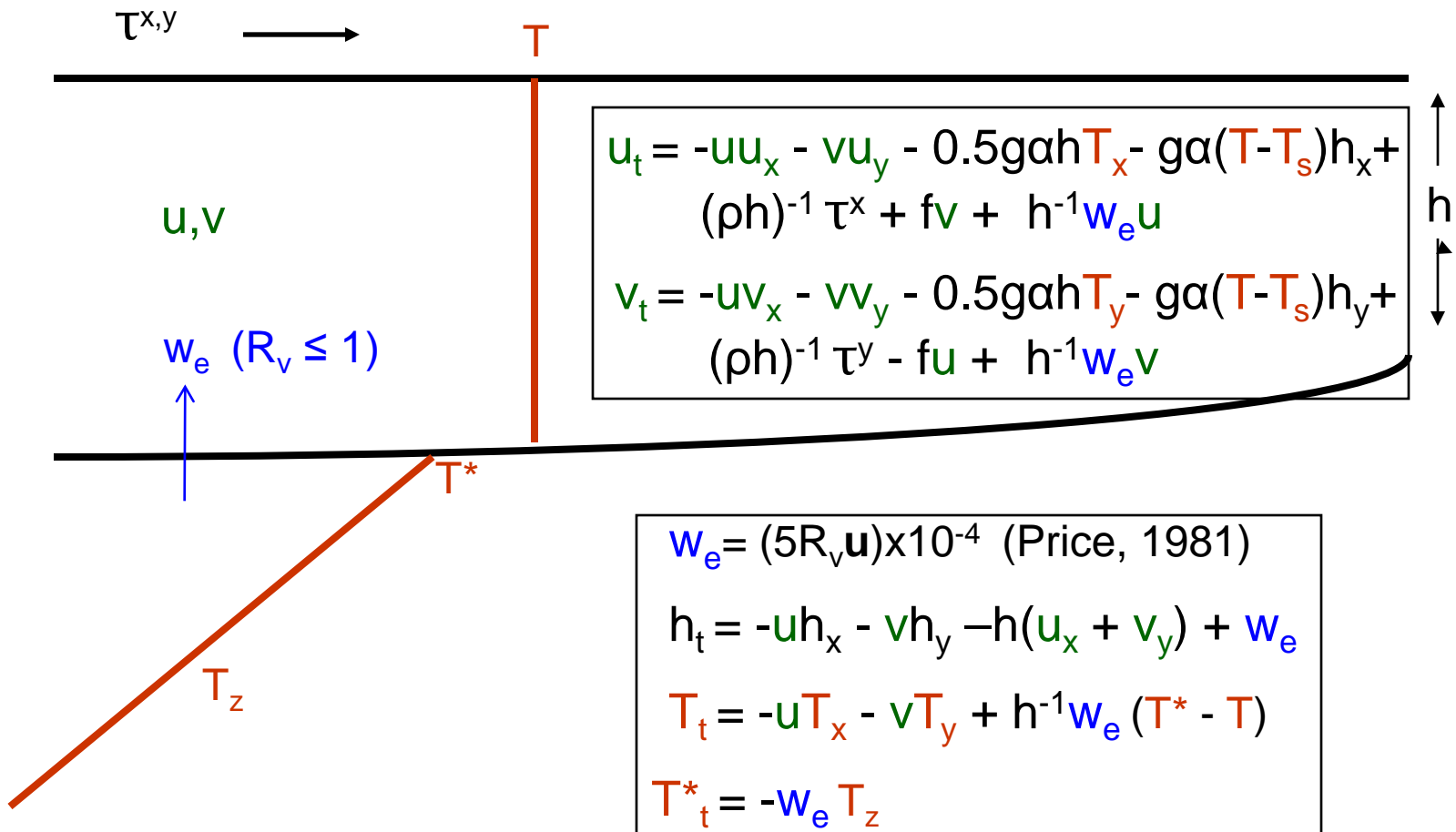
Impact on TC simulation remains to be verified –
coming soon to a theater near you....

Physics

Cumulus	None
Microphysics	Flatau (1989) and Tripoli; 2-moment prognostic scheme (specific humidity and number concentration) for all species except cloud water
PBL	K theory (horizontal), TKE (vertical)
Surface Layer	Simple Bulk ($C_H = C_M = C_D$)
Land Surface	1d soil model (Tremback and Kessler, 1985)
Radiation	RRTM (Mlawer et al., 1997a; Mlawer and Clough, 1997b)

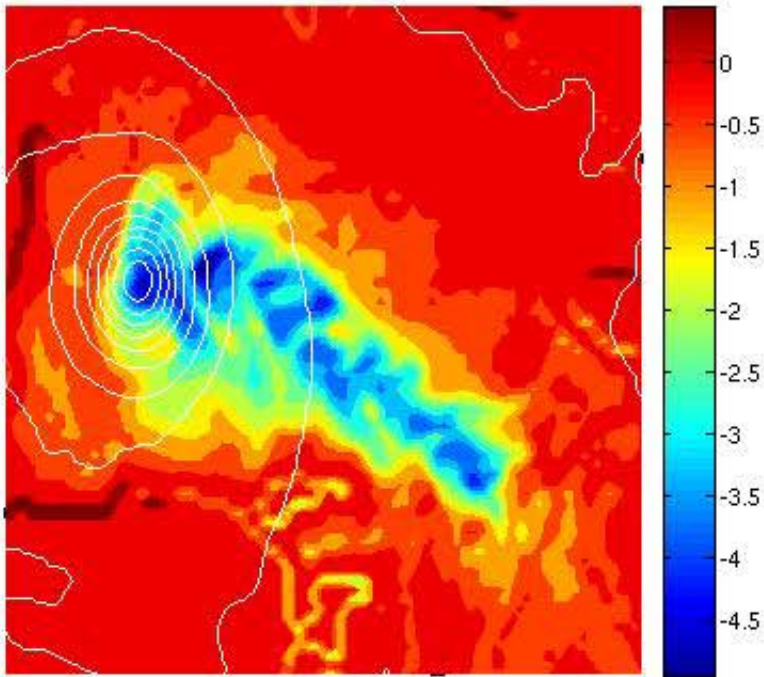
Ocean

- 1.5 – layer ocean (mixed-layer, thermocline)

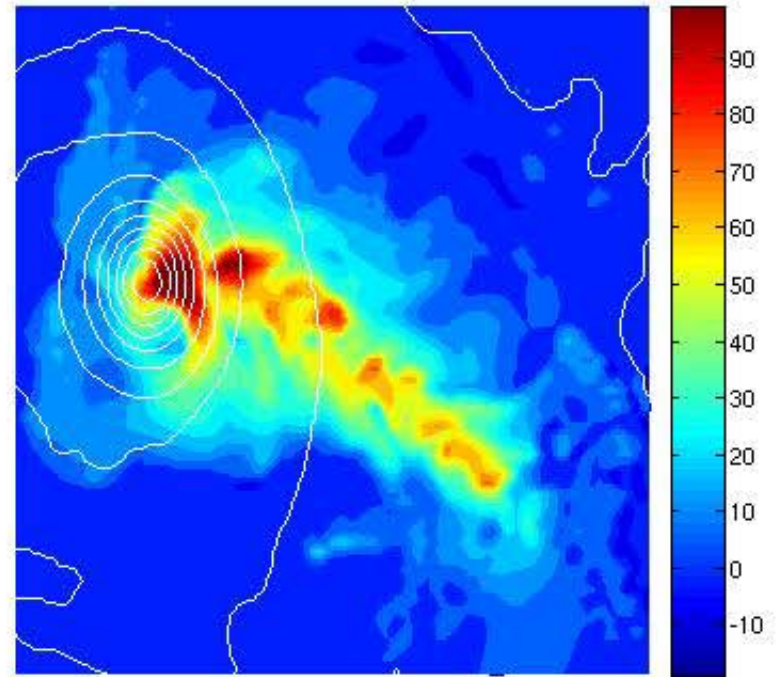


Wilma 2005101800 48-hr FCST

Δ SST ($^{\circ}$ C)



Δ h (m)



Configuration

- Atmosphere

- 12km (640x640), 3km (302x302), 1km (353x353)
- 46 levels (200m through 2km, stretched to 700m at 22.8km)
- GFDL IC, GFS BC

- Ocean

- 12km (640x640); u,v,h,T interpolated to grids 2 and 3.
- Initial SST, ML depth and subsurface temperature gradient from HYCOM north and equatorial Atlantic analyses.

Tribulations

- Porting UW-NMS Code to Bluefire turned out to be a non-trivial task
- Complete OMP implementation and optimization ~ several months
- Output module was less a problem for us (though Ligia may disagree!)

Trials

- Shared-memory environment (max 32 CPUs per run on Bluefire) BCS*:
 - 2 days for 126-hr forecast on 1 grid
 - 5 days for 126-hr forecast on 2 grids
 - 12 days for 126-hr forecast on 3 grids
- Clearly, we need (re)distributive change!

* Best Case Scenario given queue vicissitudes

Simulations

- Complete*:
 - Wilma 2005101600 (1 and 2 grids)
 - Wilma 2005101700 (1, 2 and 3 grids)
 - Wilma 2005101800 (1 and 2 grids)
- Current plan: finish the Wilma cases before 27 May (at least 3 runs at high resolution)

* Though currently being re-run

Why High Resolution?

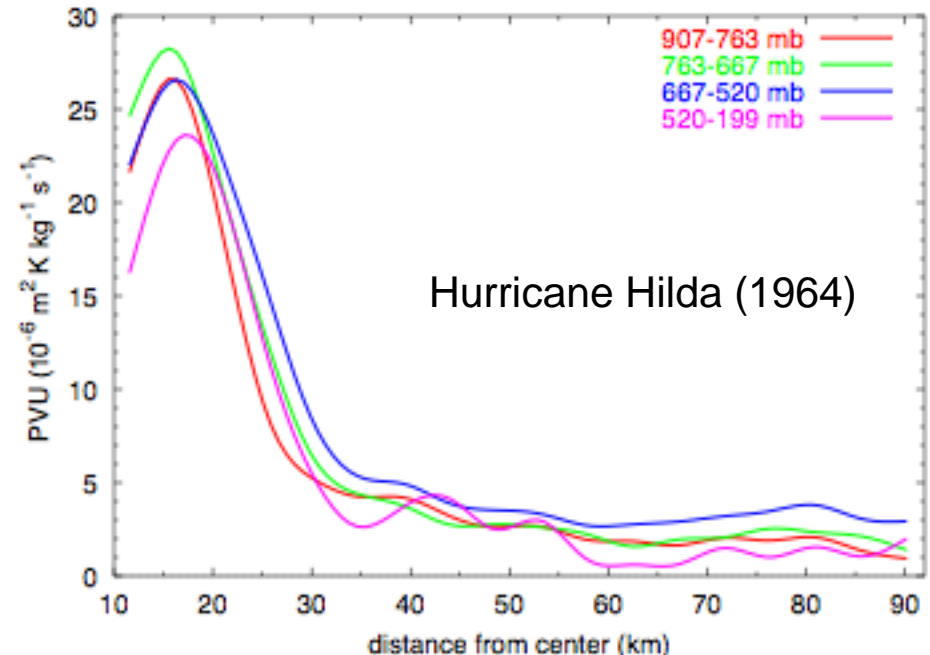
No doubt many processes / phenomena important in TC evolution benefit from an increase in resolution.

Which of these have been studied in simpler, more easily understood settings? We'll look at one.

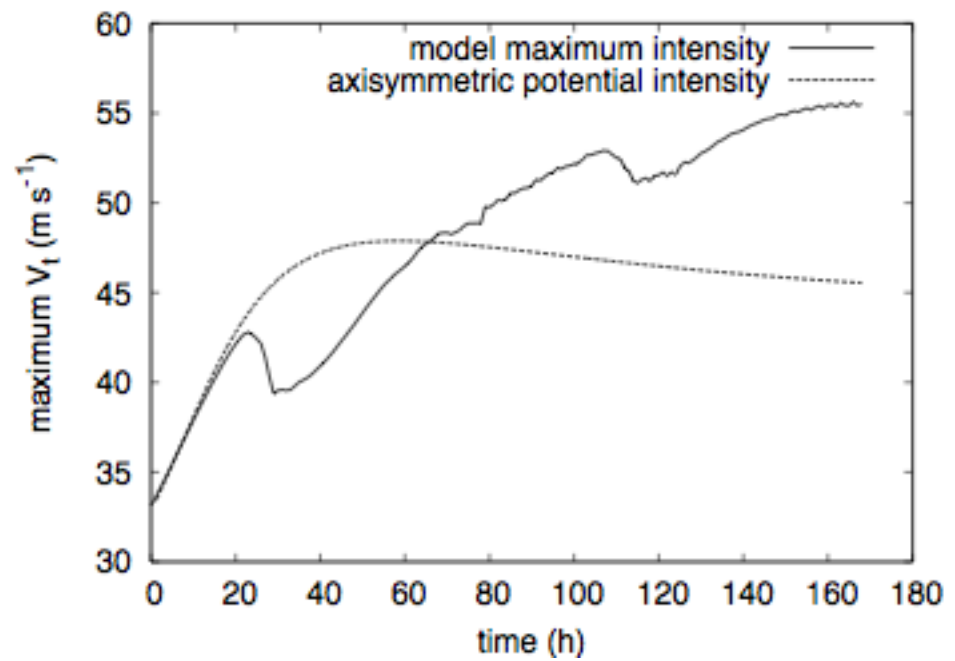
PV Mixing: application to hurricanes (Kossin et al, Rozoff et al)

As a nascent tropical storm develops, persistent diabatic heating forms a “tower” of PV. When an eye develops, diabatic heating and PV production is confined to an annulus. This results in a “hollow tower” PV structure (Möller and Smith 1994).

The radial gradient of PV changes sign and the flow can support barotropic instability (Schubert et al. 1999).

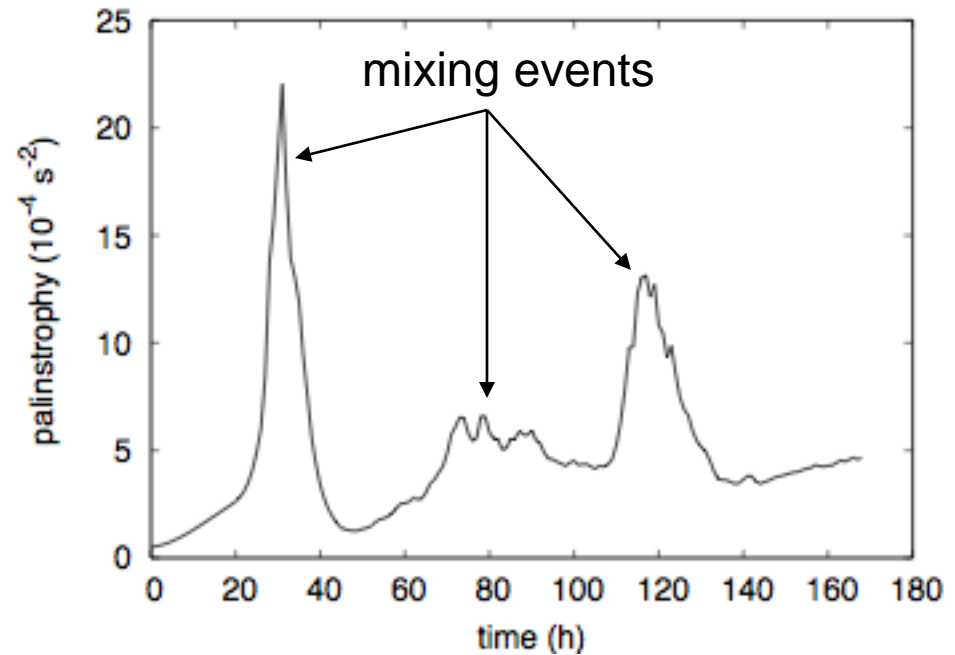


Intensity evolution



Palinstrophy evolution

$$\mathcal{P} = \iint \frac{1}{2} \nabla \zeta \cdot \nabla \zeta \, dx dy$$



Questions

Outside the idealized framework, what is the role of model resolution in the PV mixing process?

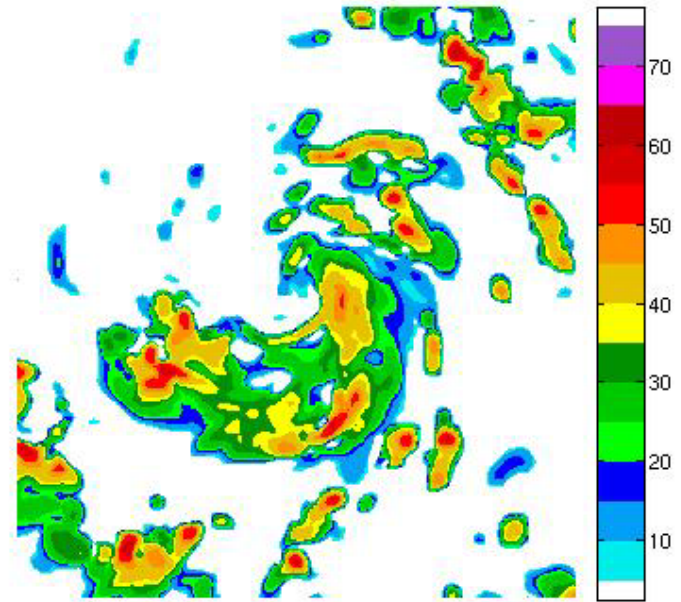
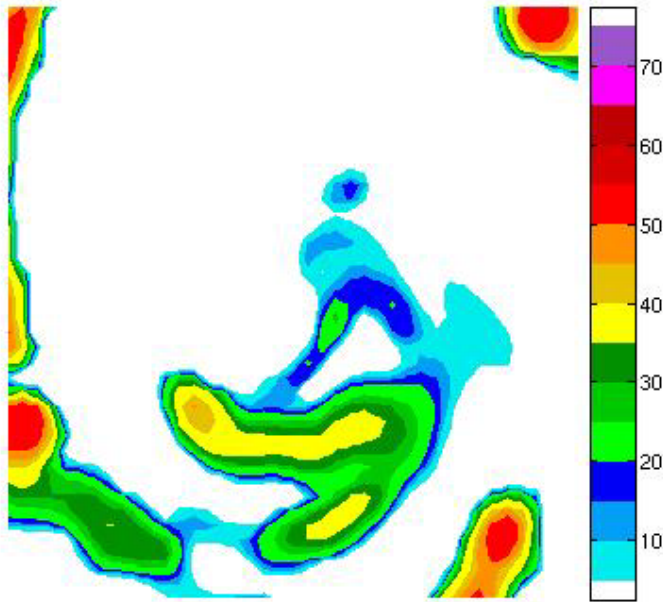
In particular, can a hurricane simulated at 12-km support the establishment of sharp (→ vortex sheet) PV gradients?

12-km

12h

3-km

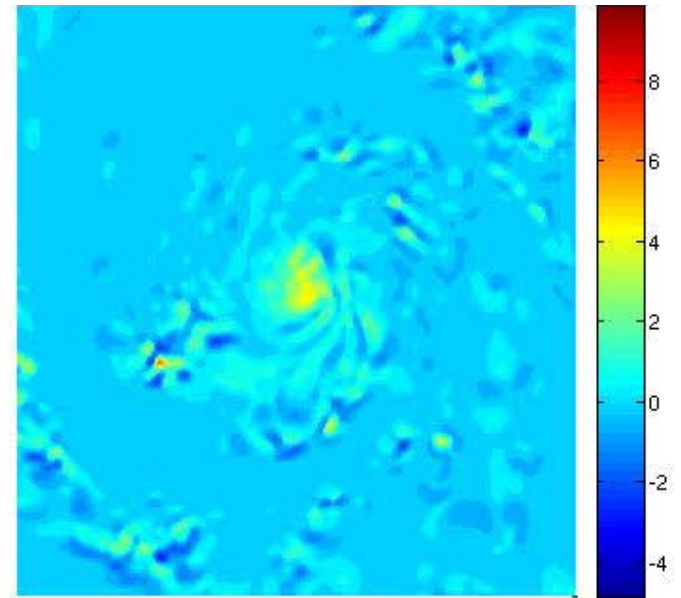
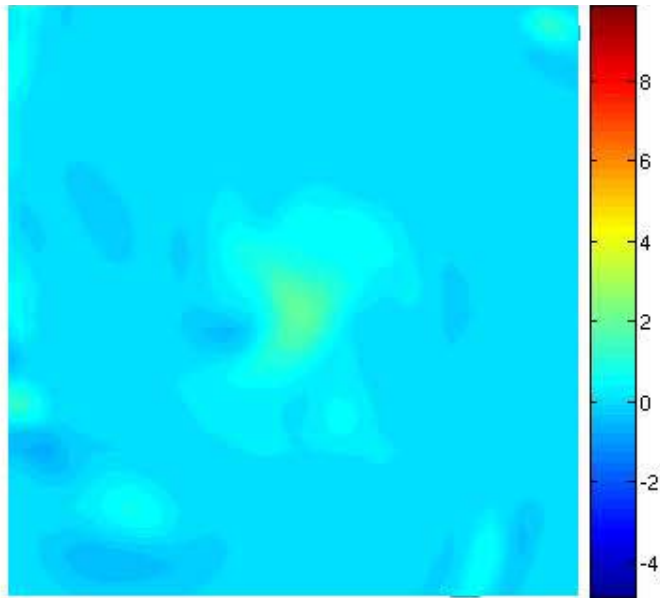
dBZ
at 2km



985 hPa, 48 kt

975 hPa, 80 kt

q_z ($\times 10^3$)
at 2km

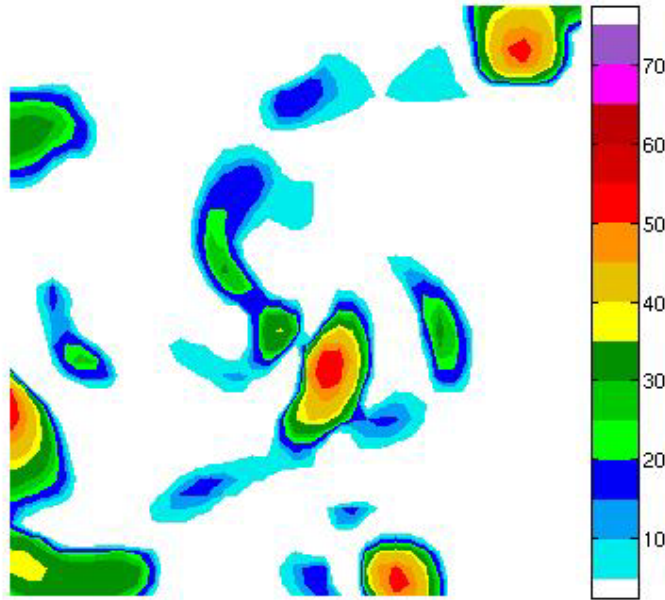


12-km

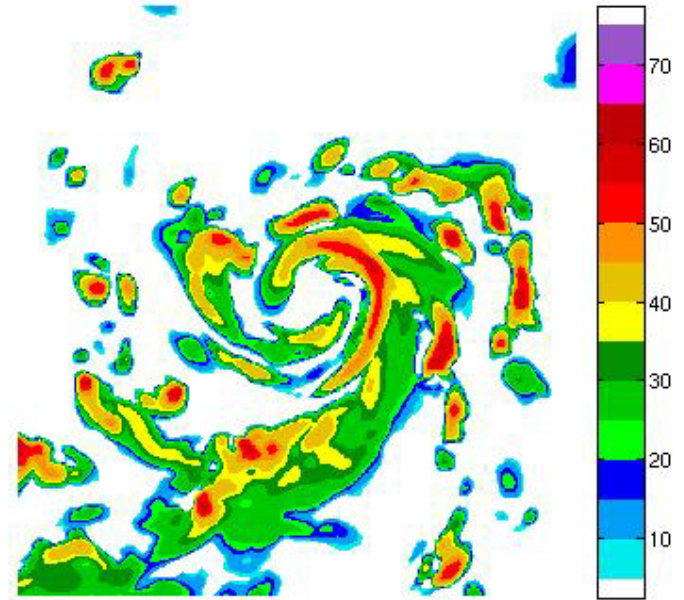
24h

3-km

dBZ
at 2km

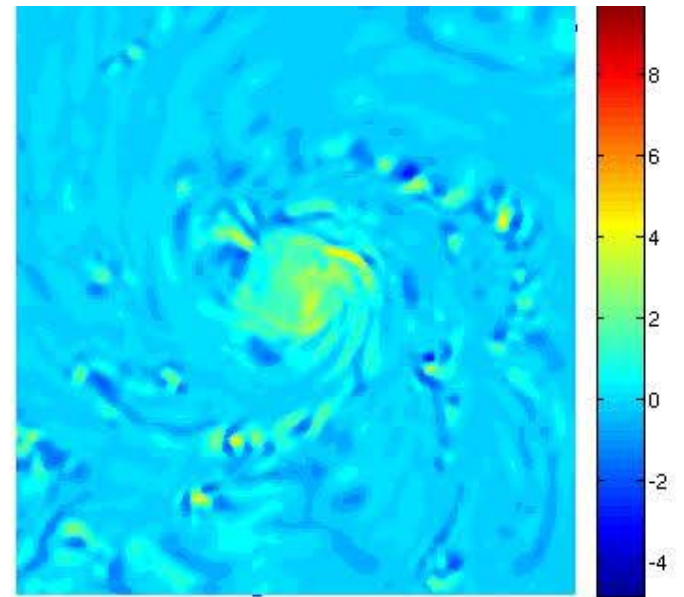
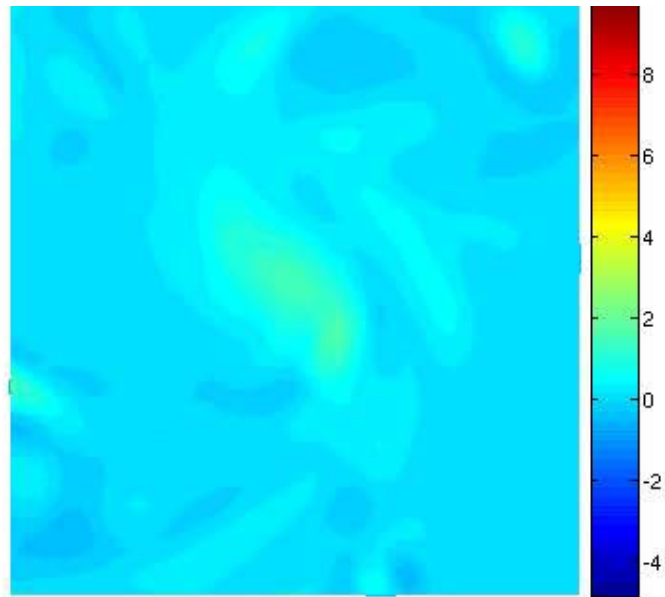


986 hPa, 50 kt



968 hPa, 88 kt

q_z ($\times 10^3$)
at 2km

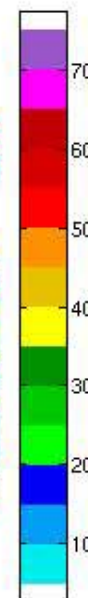
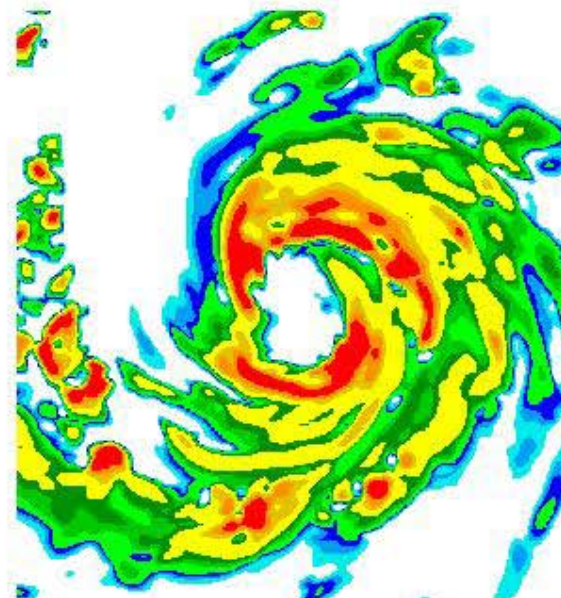
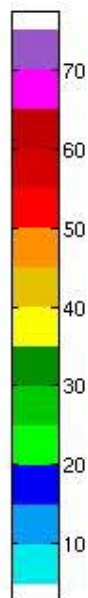
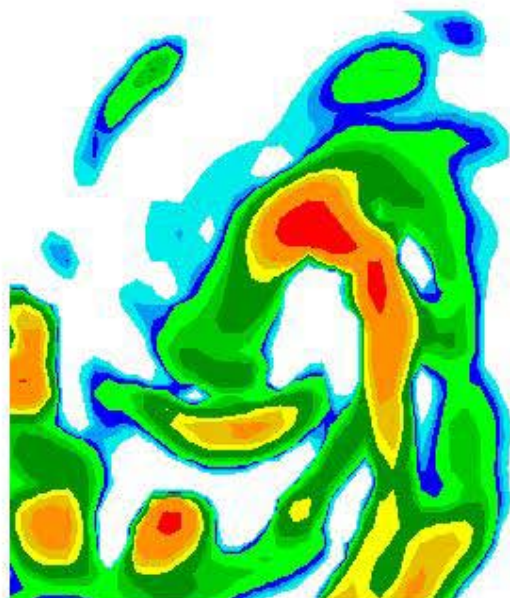


12-km

36h

3-km

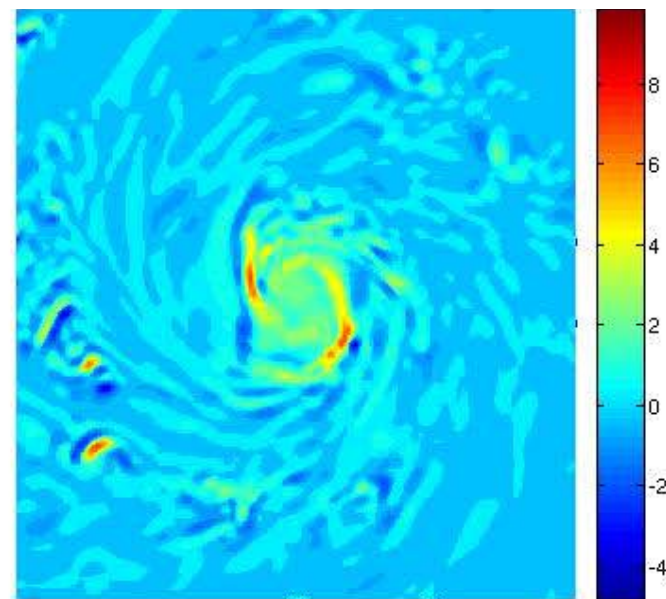
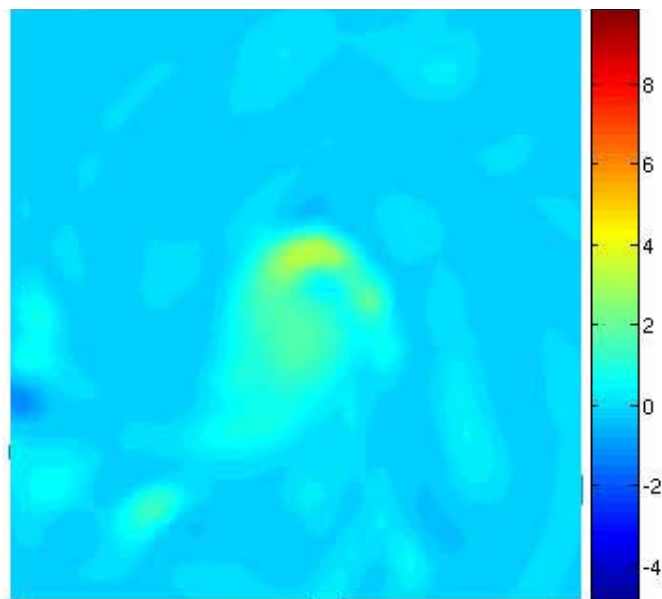
dBZ
at 2km



971 hPa, 75 kt

945 hPa, 110 kt

q_z ($\times 10^3$)
at 2km

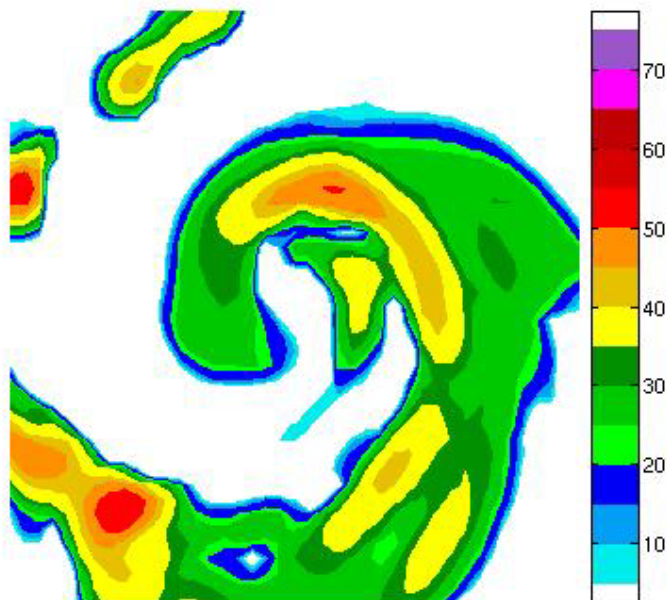


12-km

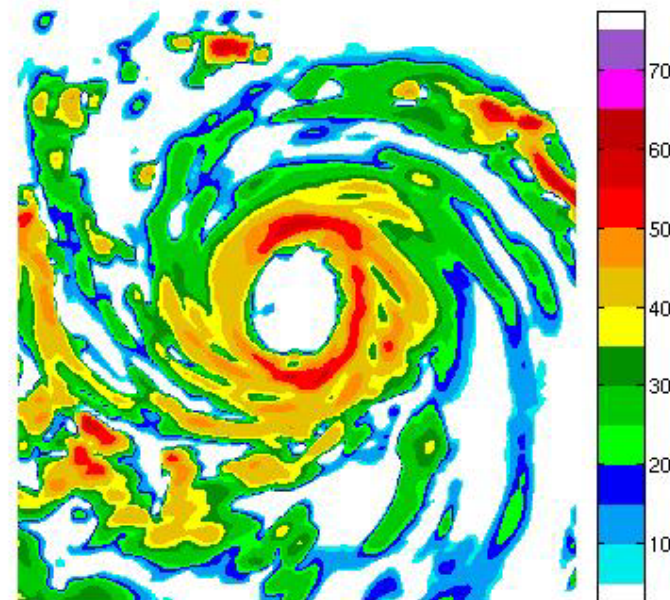
48h

3-km

dBZ
at 2km

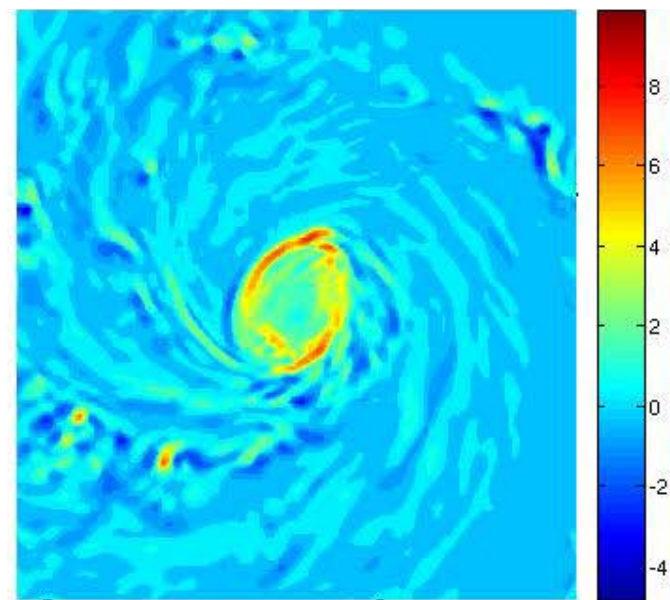
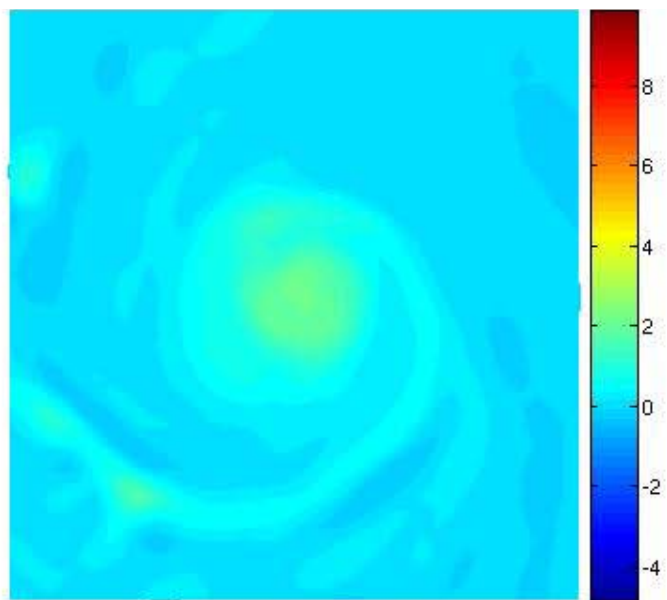


962 hPa, 80 kt



922 hPa, 132 kt

q_z ($\times 10^3$)
at 2km



Plans

- Finish analysis of HRH runs (PV-mixing episodes) at 12-, 3- and 1-km. Intensity and resolution relationships?
- Ensemble applications (group dependent):
 - Single vs. multi-model
 - Resolution vs. sample size
- EnKF (+TMI, SSM/I, GOES)

UW-NMS Upgrades

- NOAH LSM
- Sea Spray (Andreas)
- YSU PBL
- 3-D Ocean w/ wave model
- **MPI!!**
- AMPS microphysics
- PPM scalar advection